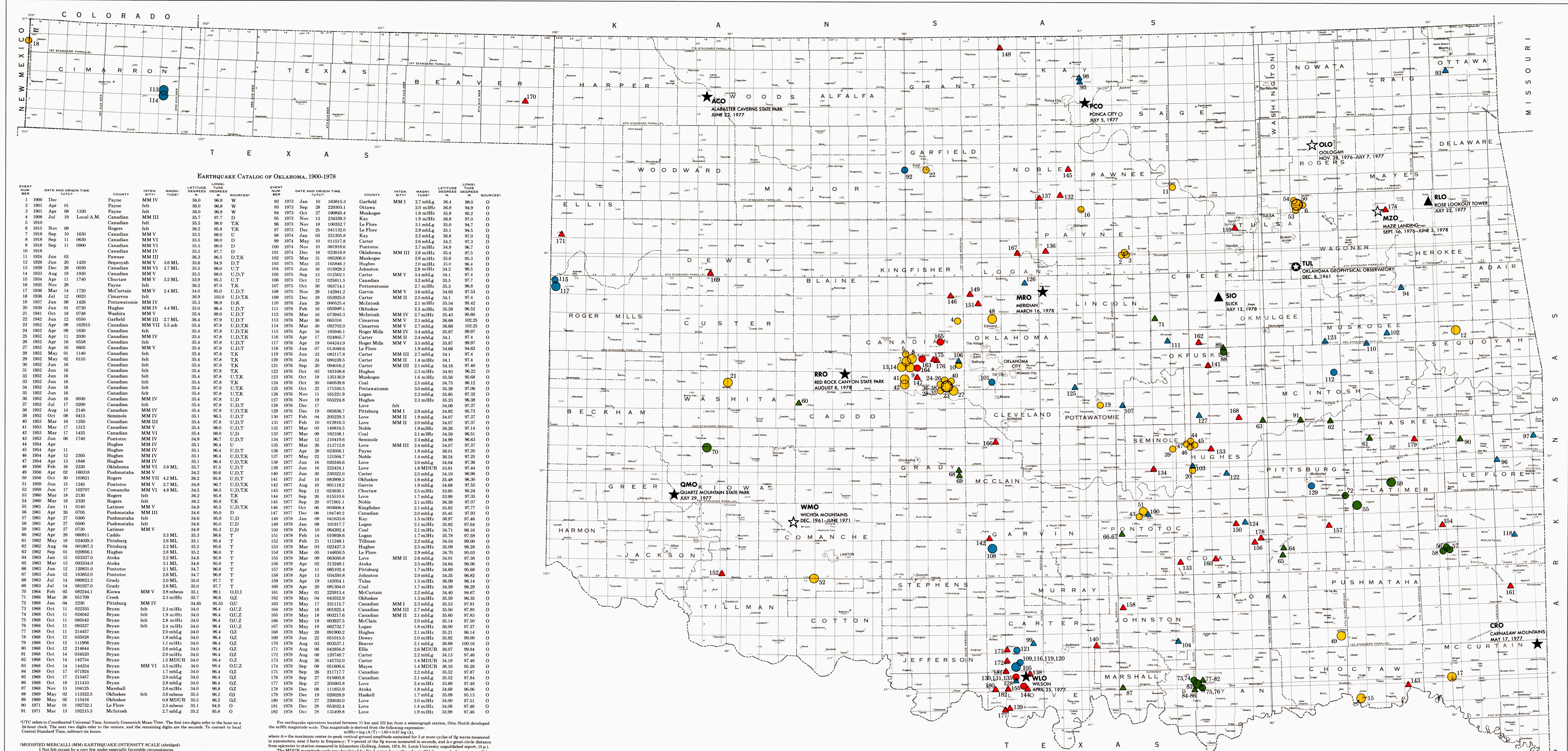




If felt only by a few persons at rest, especially on upper floors of buildings. Suspended objects may swing.
 V Felt quite noticeably by many persons, especially on upper floors of buildings. Standing motion may be weakly slight.
 IV During the day the motion is barely noticed by many, outdoors by few. At night some awaken. Dishes, doors, windows disturbed. Standing motion can rock noticeably.
 V Felt by nearly everyone, many awakened. Dishes, windows, etc., broken; unstable objects knocked over. Pendulum clocks may stop.
 VI Felt by all; many frightened and run outdoors.
 VII Everybody runs outdoors. Damage negligible in buildings of good design and construction. Shock noticed by persons driving motor cars.
 VIII Damage slight in specially designed structures; considerable in ordinary substantial buildings; great in poorly built structures. Fall of chimneys, driving motor cars disturbed.
 IX Damage considerable even in specially designed structures; well-designed frame structures thrown out of plumb. Buildings shifted foundations. Ground cracked conspicuously.
 X Well-built, sound structures destroyed. No damage to objects well fixed in the ground. Landslides and shifting of sand dunes noted.
 XI Few if any (masonry) structures remain standing. Broad fissures in ground.
 XII Damage total. Walls and ground surface on ground surface.

Modified Mercalli Intensity Scale of 1901; Wood, H. A., and Newman, F., 1931, *Bulletin, Seismological Society of America*, vol. 21, no. 4, p. 277-283.

*Earthquake magnitude is a measurement of energy and is based on data from seismograph records. There are several different scales used in the catalog to report magnitude. They are as follows: M_L (Richter), m_b (Gutenberg and Richter), m_s (Nuttall), M_{sc} (Nuttall), and M_{DE} (Lawson). Each magnitude scale was established to accommodate specific criteria, such as the distance to the epicenter, as well as the availability of certain seismic data.

The Richter magnitude, M_L , as the amplitude, is derived from the following equation:

$$M_L = \log A - \log A_0$$
 where A = the maximum center-to-peak amplitude of Sg wave measured, in millimeters, on a seismogram recorded on short-period, vertical seismograph with a frequency recorded that does not differ greater than 1% from the design frequency;
 A_0 = displacement magnification of the vertical-motion seismograph at periods below 0.8 second; $-\log A_0$ = Richter's empirical distance correction function for southern California; and Δ = $\log$ of the great-circle distance from epicenter to seismograph measured in kilometers. This equation is Eystein Yrgeysson's unpublished adaptation to Charles Richter's magnitude scale for southern California, when applied to Oklahoma (Richter, C. F., 1958, *Earthquake seismology*; W. H. Freeman, San Francisco, p. 340-345).

The Gutenberg and Richter magnitude, m_b , is defined as

$$m_b = \log (A/T) + 1.67 + 0.007 \Delta$$
 where A = the maximum center-to-peak vertical amplitude of P wave measured in microbars, recorded at distances of 3 and 100 great-circle degrees from the epicenter; T = dominant P wave period measured in seconds; Δ = Gutenberg and Richter's empirical distance and depth correction factor (Richter, C. F., 1958, *Elementary seismology*; W. H. Freeman, San Francisco, p. 347-360, 660).

Nuttall's earthquake magnitude, m_s , for seismograph stations located between 55.6 km and 445 km from the epicenter, is derived from the following equation:

$$m_s = \log (A/T) - 1.10 + 0.09 \log \Delta$$
 When seismograph stations are located between 445 km and 3,900 km from the epicenter, m_s is defined as

$$m_s = \log (A/T) - 3.71 + 1.06 \log \Delta$$
 where A = the maximum center-to-peak vertical ground amplitude sustained for 3 or more cycles of Sg wave measured in nanometers; T = period of the Sg wave measured in seconds; and Δ = the great-circle distance from station to epicenter measured in kilometers (Nuttall, O. W., 1973, *Seismic wave attenuation and magnitude relations for eastern North America*; Journal of Geophysical Research, v. 78, no. 5, p. 876-885).

The Woodens magnitude scale, m_b , is defined as

$$m_b = \log (A/T) + 3.27 + 2 \log \Delta$$
 where A = the center-to-peak vertical ground amplitude of P₀ waves, measured in nanometers, recorded between 200 km and 2,100 km from the epicenter; T = period of the P₀ wave measured in seconds; and Δ = great-circle distance in kilometers (Woodward, J. E., 1967, *Empirical determination of epicenter and near-shear-slip rate from the seismicity of the United States*; Bulletin of the Seismological Society of America, v. 57, no. 4, p. 901-928).

The magnitude scale was developed by Sir Lawson for earthquakes in Oklahoma and adjacent areas. It is defined as

$$M_{DE} = 1.86 \log (DUR) - 1.67$$
 where DUR = the duration or difference, in seconds, between $T = 1.87$ s and the time the final coda amplitude is measured, or twice the background noise amplitude. If the P₀ wave is the first arrival, the interval between the coda amplitude and the P₀ wave is used for the duration. The duration is measured in seconds (Lawson, J. E., Jr., 1978, A preliminary duration magnitude scale for local and regional earthquakes recorded at Oklahoma Seismograph Stations; Oklahoma Geophysical Observatory, open file report, 14 p.).

Sources for Earthquake Data for the Catalog:
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 K: Kell, B.H., 1964, Earthquakes that shook Oklahoma: Orit's Magazine of The Sunday Oklahomaian, Oklahoma City, September 27, p. 47.
 T: Yrgeysson, Eystein, 1964, Seismicity of Oklahoma: Seismological Society of America annual meeting, unpublished presentation.
 U: United States Earthquake annual publications of U. S. Department of Commerce, Coast and Geodetic Survey, 1929 through 1966; National Oceanic and Atmospheric Administration (NOAA), National Ocean Survey for 1969; NOAA, Environmental Data Service Area for 1970 through 1972; and U. S. Department of Commerce, and U. S. Department of Interior, Geological Survey for 1973 to present.
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 I: International Seismological Centre, Edinburgh, Scotland, 1964, Bulletin, v. 1, no. 2, p. 9 and 5, p. 24, 1969. (Redacted to November, England in September 1973).
 W: Woods, Laura Loos, 1975, Young Cretaceous in Oklahoma Territory: Frontier Perspectives, Inc., Stillwater, Oklahoma, 221 p.

EARTHQUAKE MAP OF OKLAHOMA

Compiled by

James E. Lawson, Jr., Robert L. DuBois, Paul H. Foster, and Kenneth V. Luza

mented earthquake in Oklahoma occurred near Cushing in December 1900. This event was followed by two additional earthquakes in the same area in April 1901.

The earthquakes in this data are grouped into four time-period intervals. The first time period, 1900 to 1960, represents an interval in which all earthquakes in Oklahoma were either located from historical accounts or from seismograph stations outside the state. In late 1961, the first two seismograph stations were installed in Oklahoma. One was located in the Wichita Mountains, WMO by the United States Air Force, and the other was built near Leonard, TUL, by the Energy Production Research Co., which operated it until 1965. From 1965 to 1978, TUL was operated by the University of Oklahoma. In 1978, it was transferred to the Oklahoma Geological Survey and named the Oklahoma Geophysical Observatory. The short-period seismograms at these stations utilized passbands, between 0.5 to 1.5 hertz, which provided good records for distant earthquakes. Although the location and detection of small local earthquakes were greatly improved by the presence of these two stations, the records from seismographs with short-period passbands are of little use in recording small local earthquakes. In 1973, the Oklahoma Geo-

physical Observatory (TUL) began to record selected high-frequency seismic responses at increased magnifications. This greatly increased our capability to detect small earthquakes and to discriminate earthquakes from quarry blasts. However, the location of earthquakes depended either on finding felt reports and/or having clear first-motion directions and horizontal-amplitude measurements made from the seismic records at the Oklahoma Geophysical Observatory. In 1976, a program was begun to install seven semipermanent, volunteer-operated seismograph stations and three radio-telemetry stations in Oklahoma with support from the U.S. Nuclear Regulatory Commission. These additional seismograph stations have tremendously improved earthquake detection and location accuracy in the State of Oklahoma as well as in the adjacent states.

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HADLEY, J. B., and DEVINE, J. F., 1974, Seismotectonic map of the eastern United States: U.S. Geological Survey Miscellaneous Field Study Map MF-620, 8 p., 3 sheets.

KING, P. B., 1961, The tectonics of middle North America—east of the Cordillera system: Princeton University Press, Princeton, N.J., 293 p.

It is intended for use as a guide only to earthquake intensity and epicentral locations. Epicentral locations are based on data that may vary greatly in accuracy, particularly earthquakes that occurred prior to 1960.

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HADLEY, B., and DEWINE, J. F., 1974, Seismotectonic map of the eastern United States, U.S. Geological Survey Miscellaneous Field Study Map MF-620, 9, 3 sheets.
KOPPEL, F. H., 1960, The tectonic and seismic America-east of the Canadian system, Princeton University Press, Princeton, N.J., 289 p.

to event number in earthquake catalog

Intensity values during selected time period

Seismoperimeter field station, active; includes station abbreviation, name, and installation date

Seismoperimeter field station, inactive; includes station abbreviation, name, and operating date

Radio telemetry field station, includes station abbreviation, name, and installation date

Oklahoma Geophysical Observatory